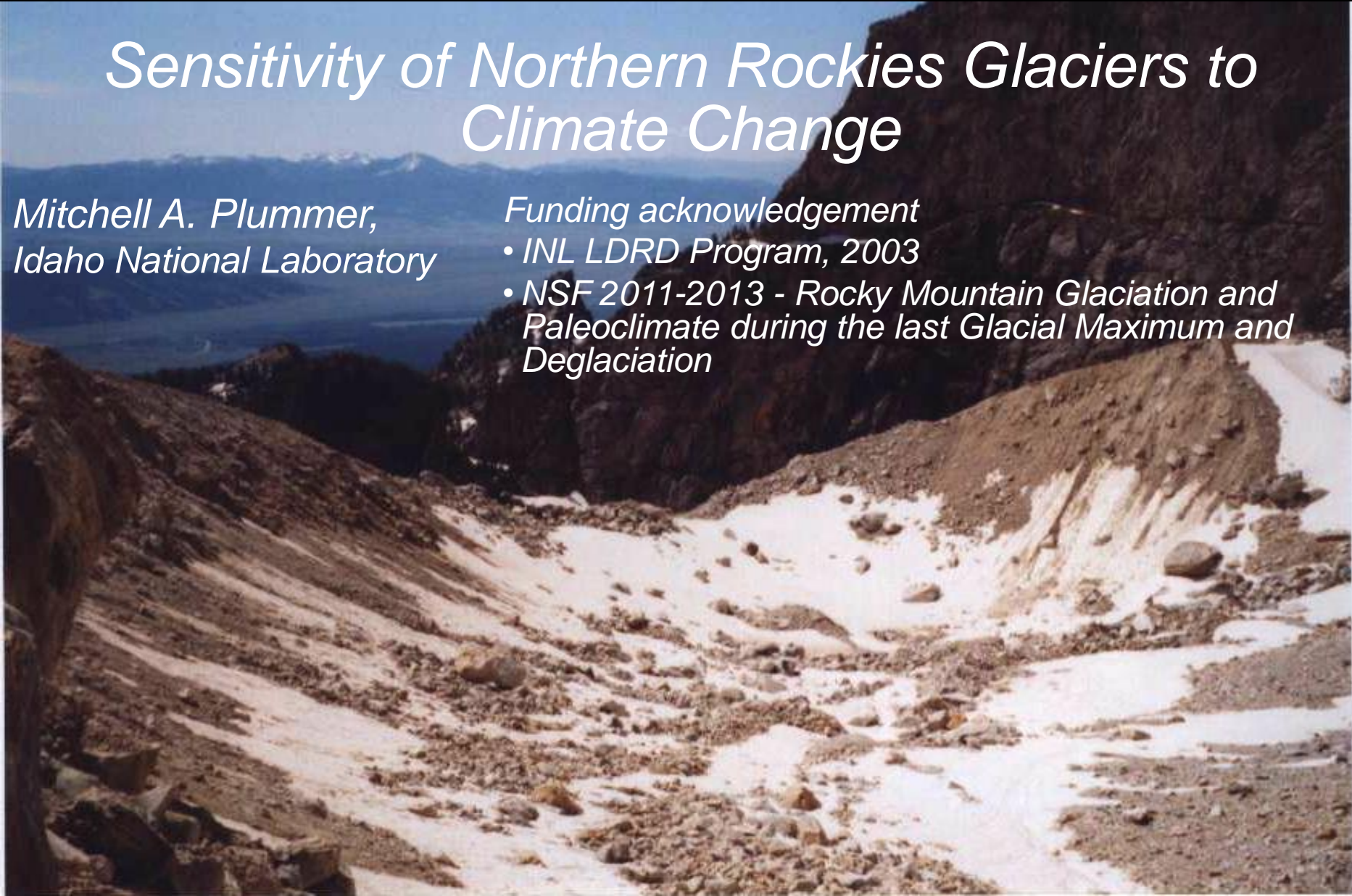


Sensitivity of Northern Rockies Glaciers to Climate Change

*Mitchell A. Plummer,
Idaho National Laboratory*

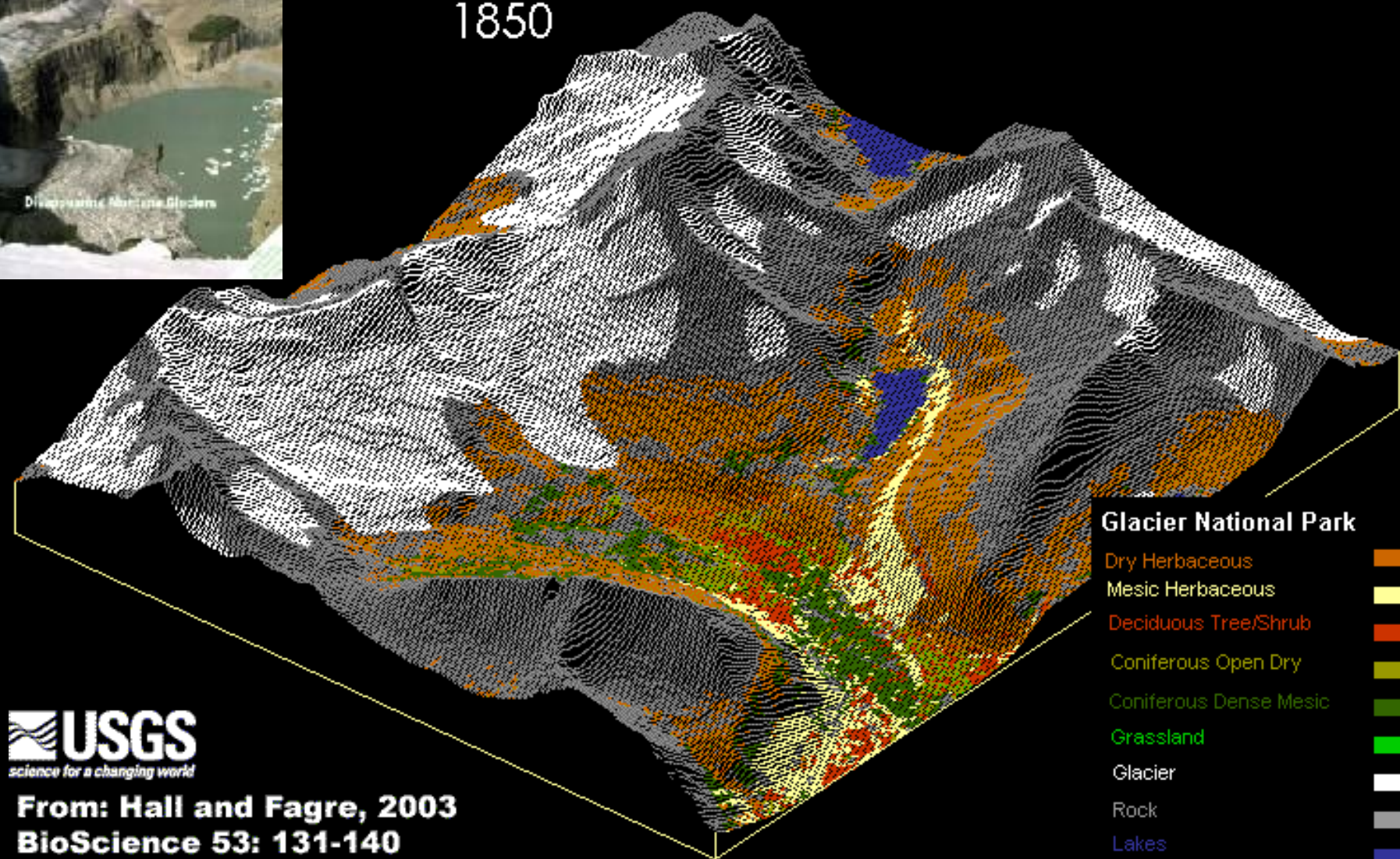
Funding acknowledgement

- *INL LDRD Program, 2003*
- *NSF 2011-2013 - Rocky Mountain Glaciation and Paleoclimate during the last Glacial Maximum and Deglaciation*



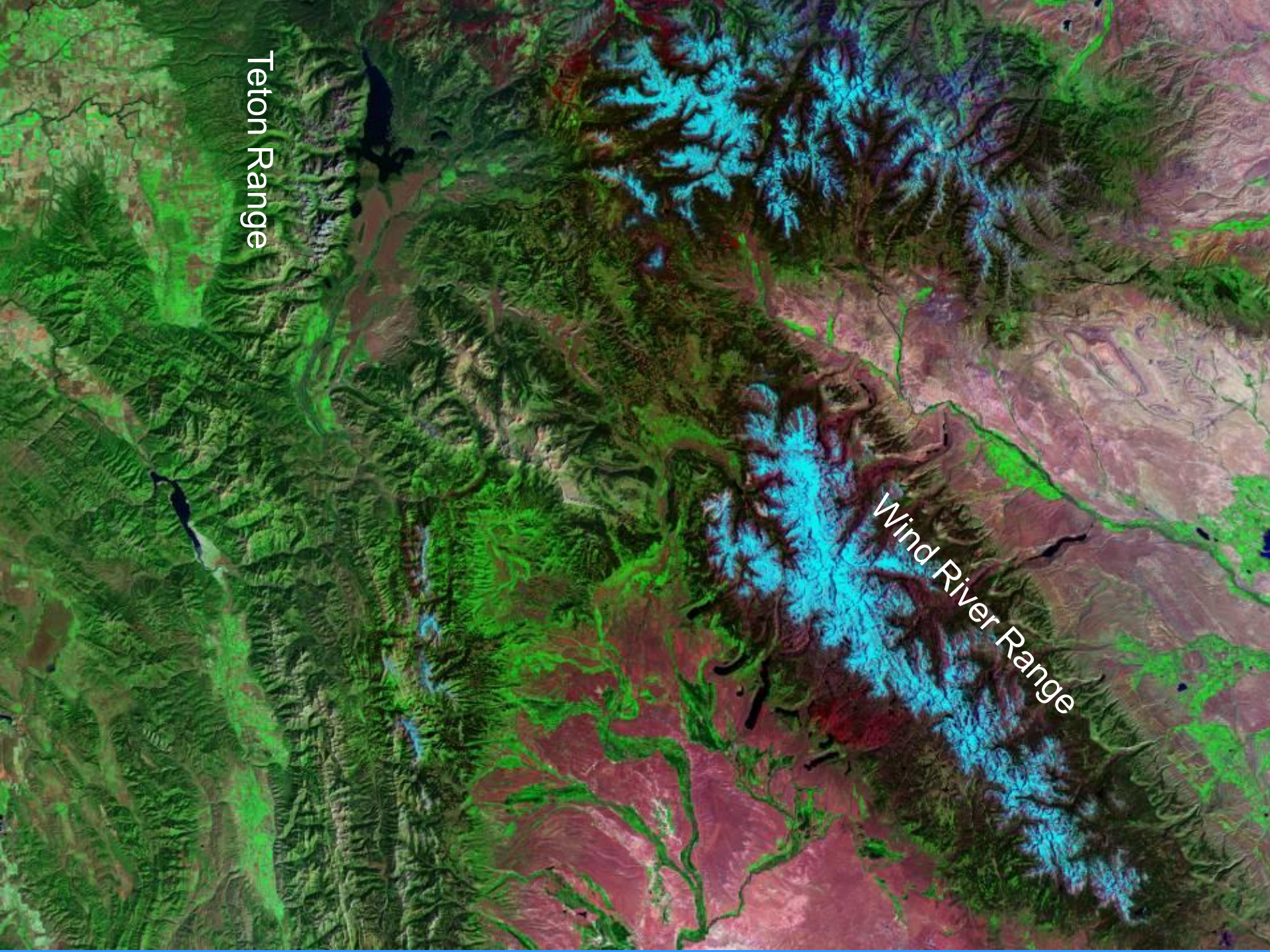
1850

Dispersing Montana Glaciers



Wind River Range

Teton Range

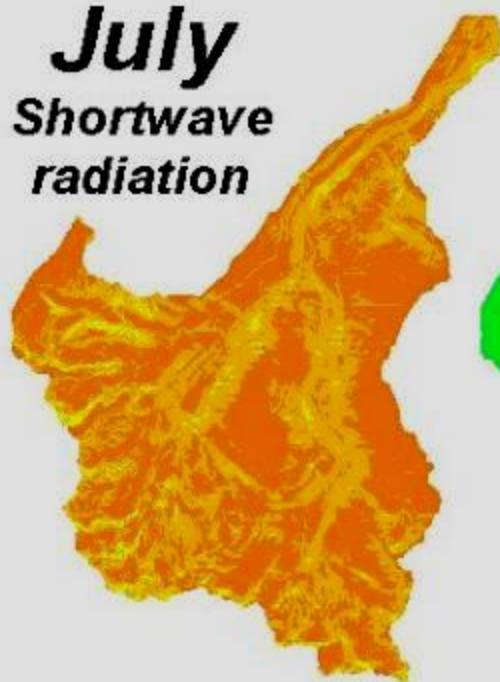


Net annual ice balance, b_n

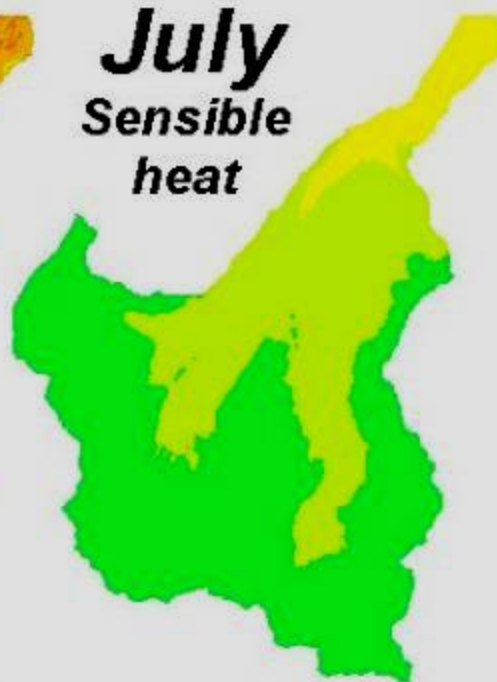
$$b_n = \int_{fall}^{spring} (Snow - E)dt + \int_{spring}^{fall} (Snow - M - E)dt$$

- b_n = net ice budget at any location in the x,y plane
- Snow = total snowfall
 - Fraction of precipitation falling as snow
- M = mass of snow melted
 - Calculated from energy balance when $T > 0$
- E = evaporation & sublimation
 - Calculated via turbulent heat transfer equations

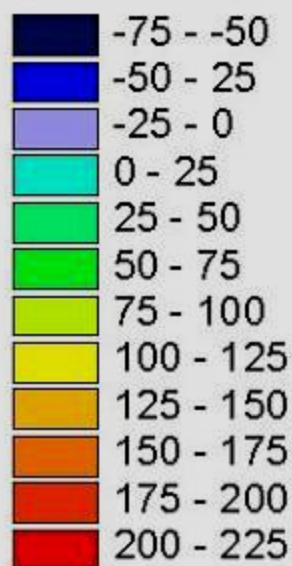
July
Shortwave
radiation



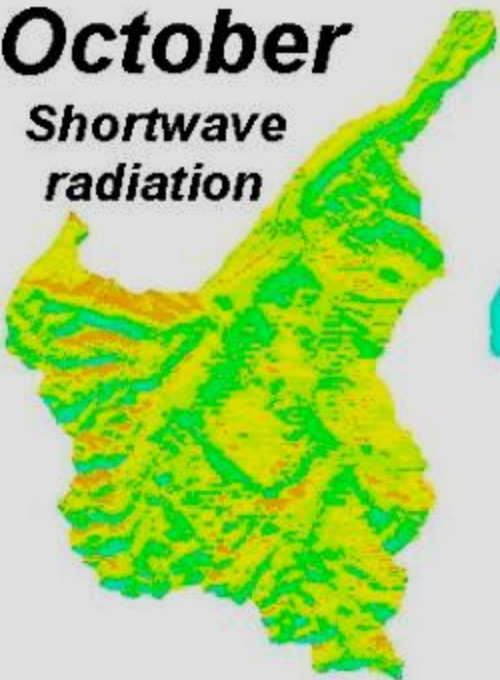
July
Sensible
heat



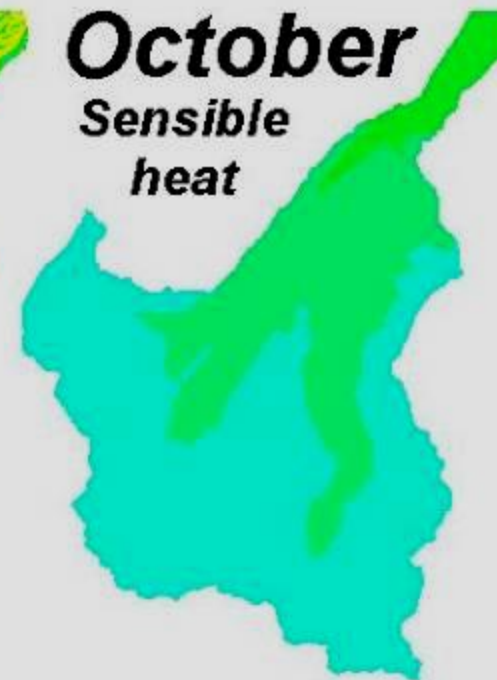
Monthly
Energy
Input
(watts/m2)



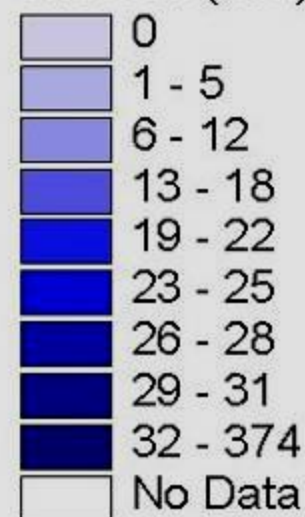
October
Shortwave
radiation



October
Sensible
heat

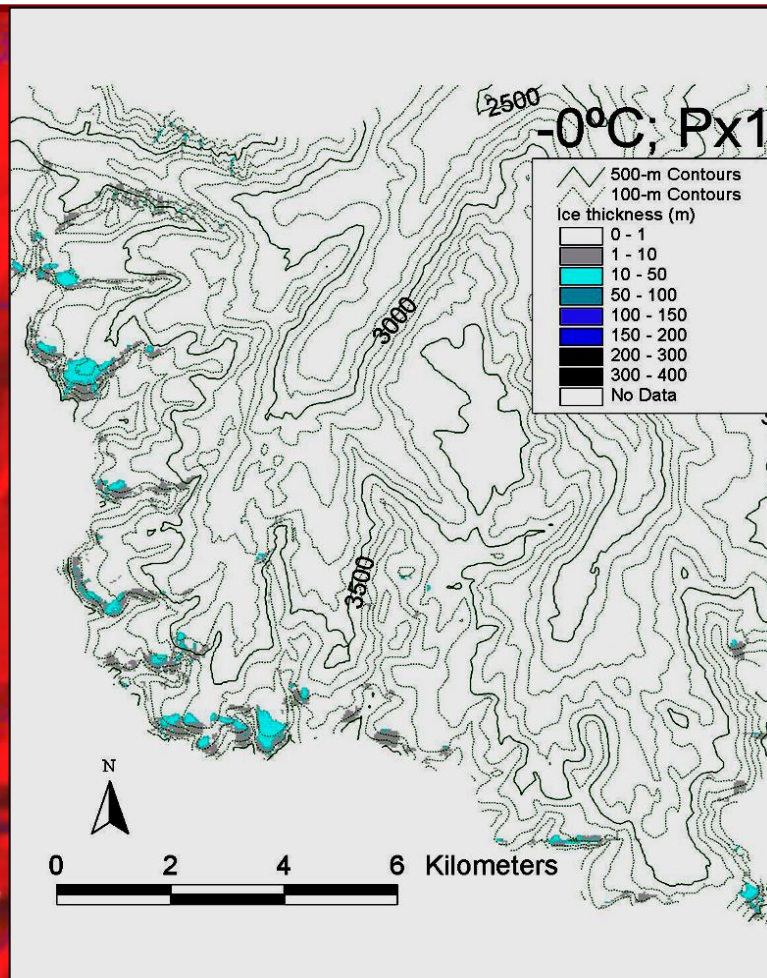
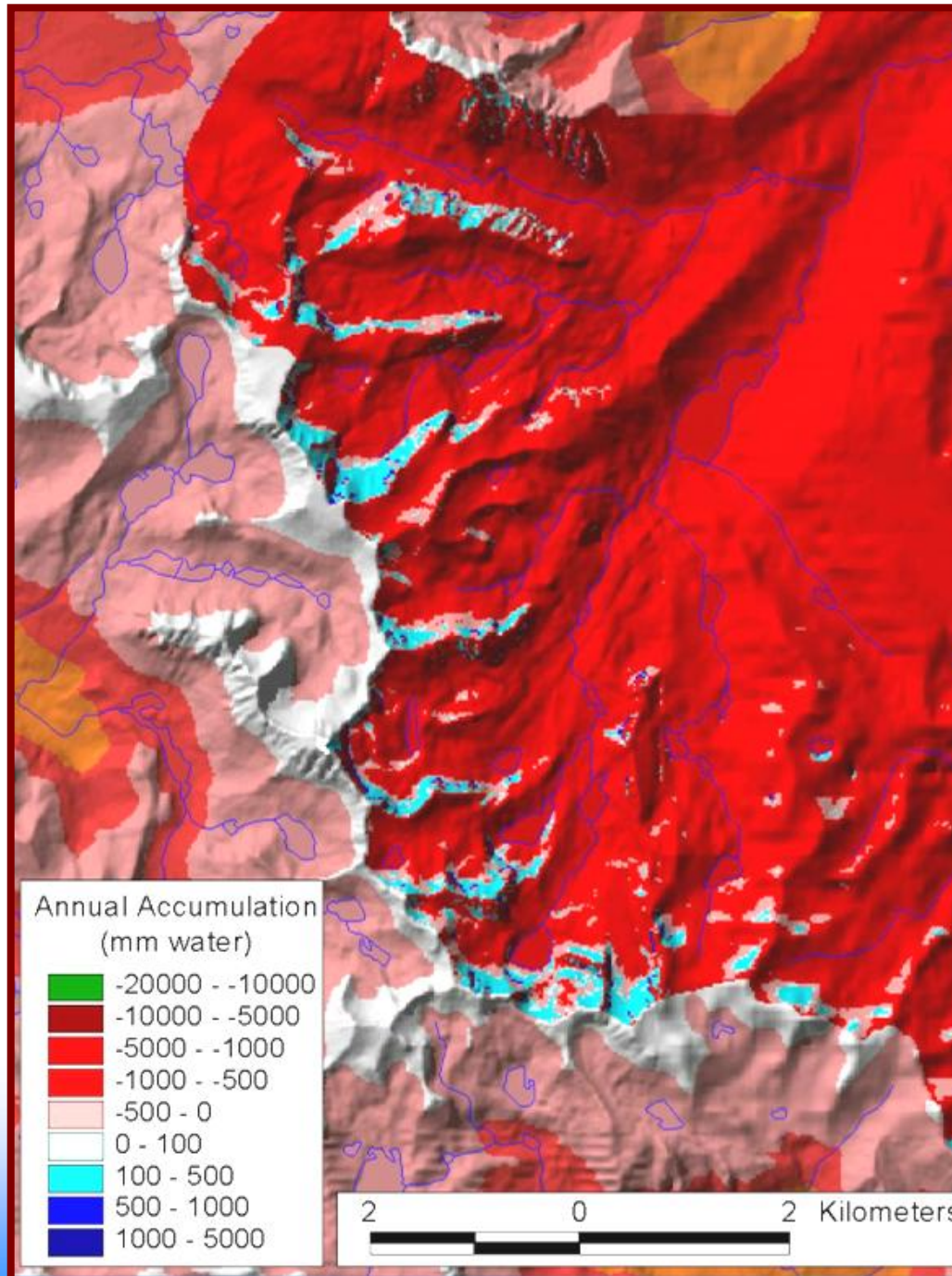


Snowfall (mm)

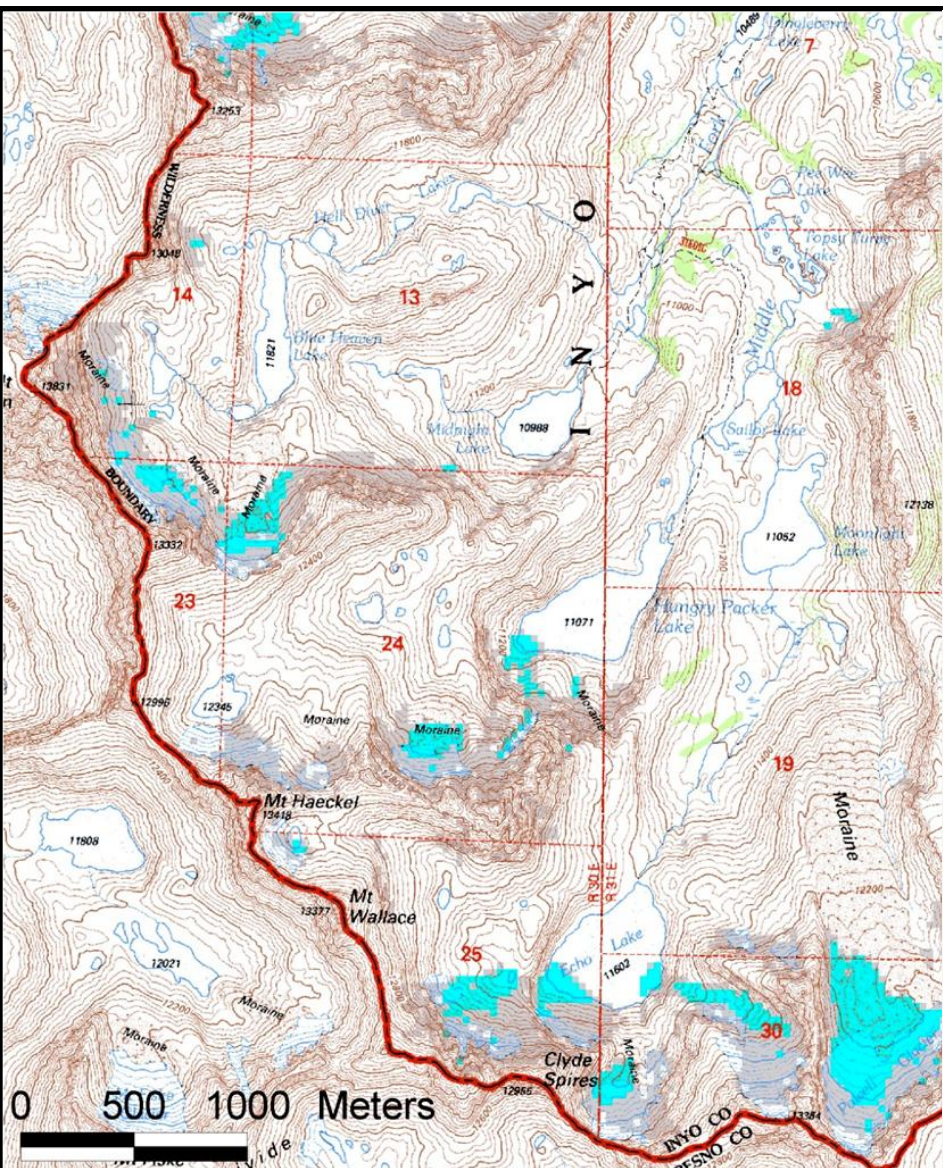


October
Snowfall

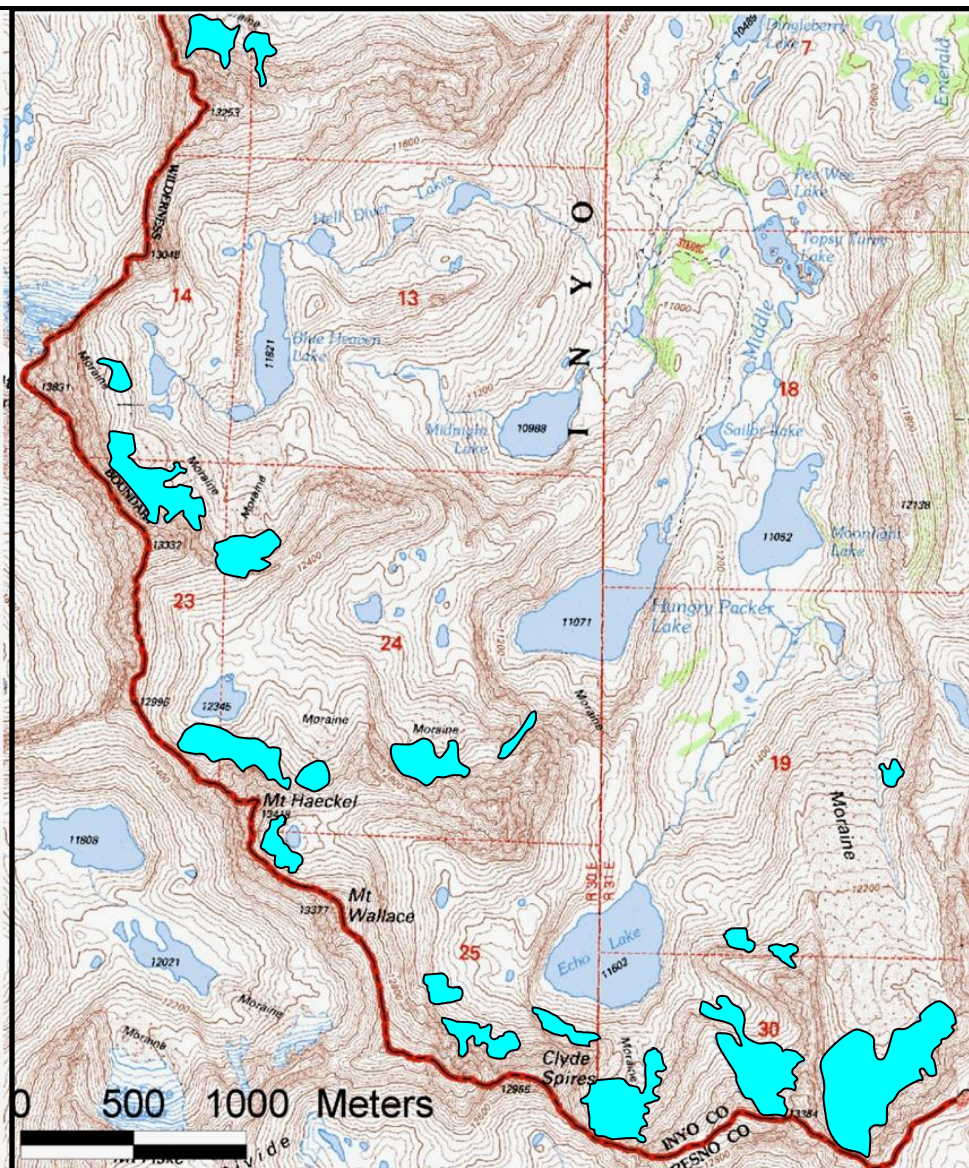




Modeled Glaciers and Perennial Snow

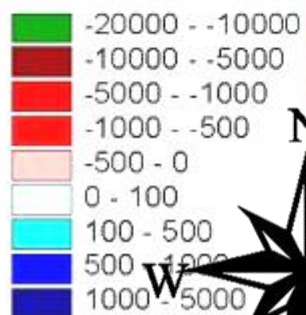


Mapped Glaciers and Perennial Snow



$\Delta T = -6^{\circ}\text{C}$

Annual Accumulation
(mm water)



N

W

E

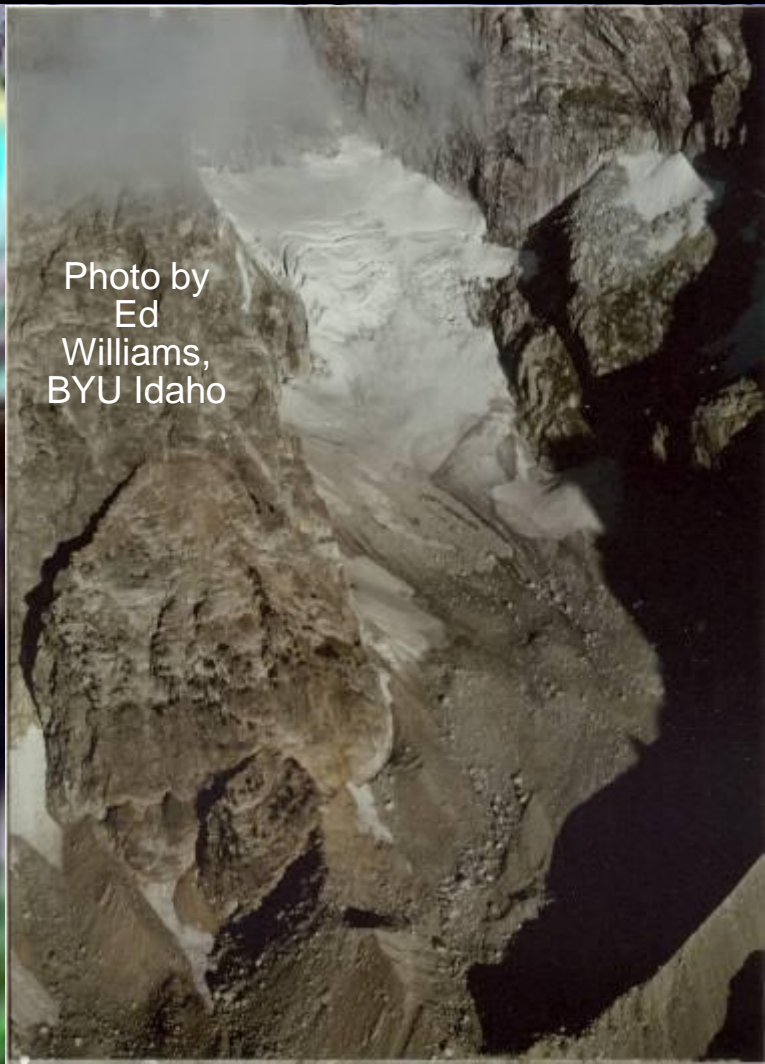
0

2

Kilometers



Photo by
Ed
Williams,
BYU Idaho



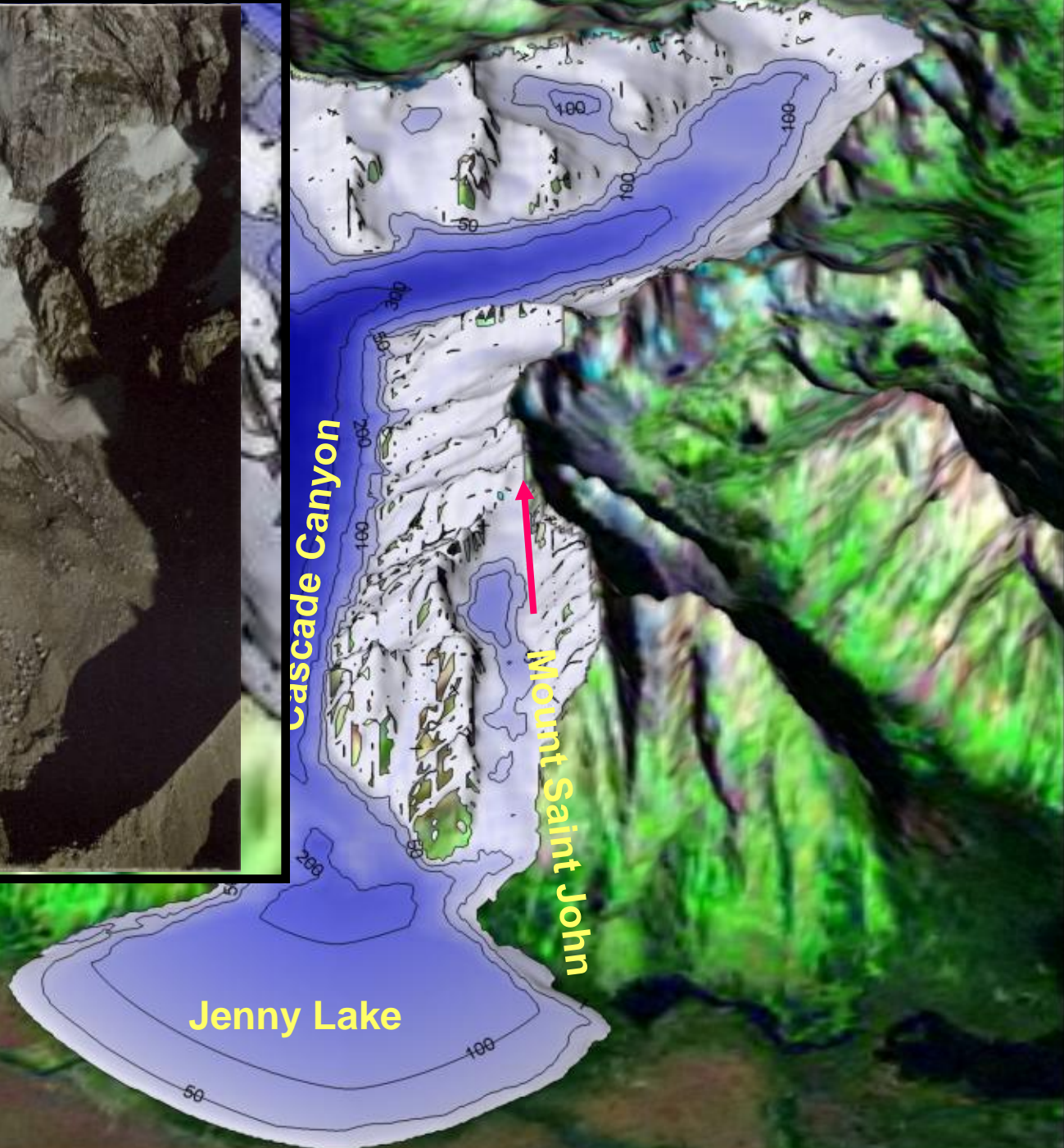
Ice thickness (m)



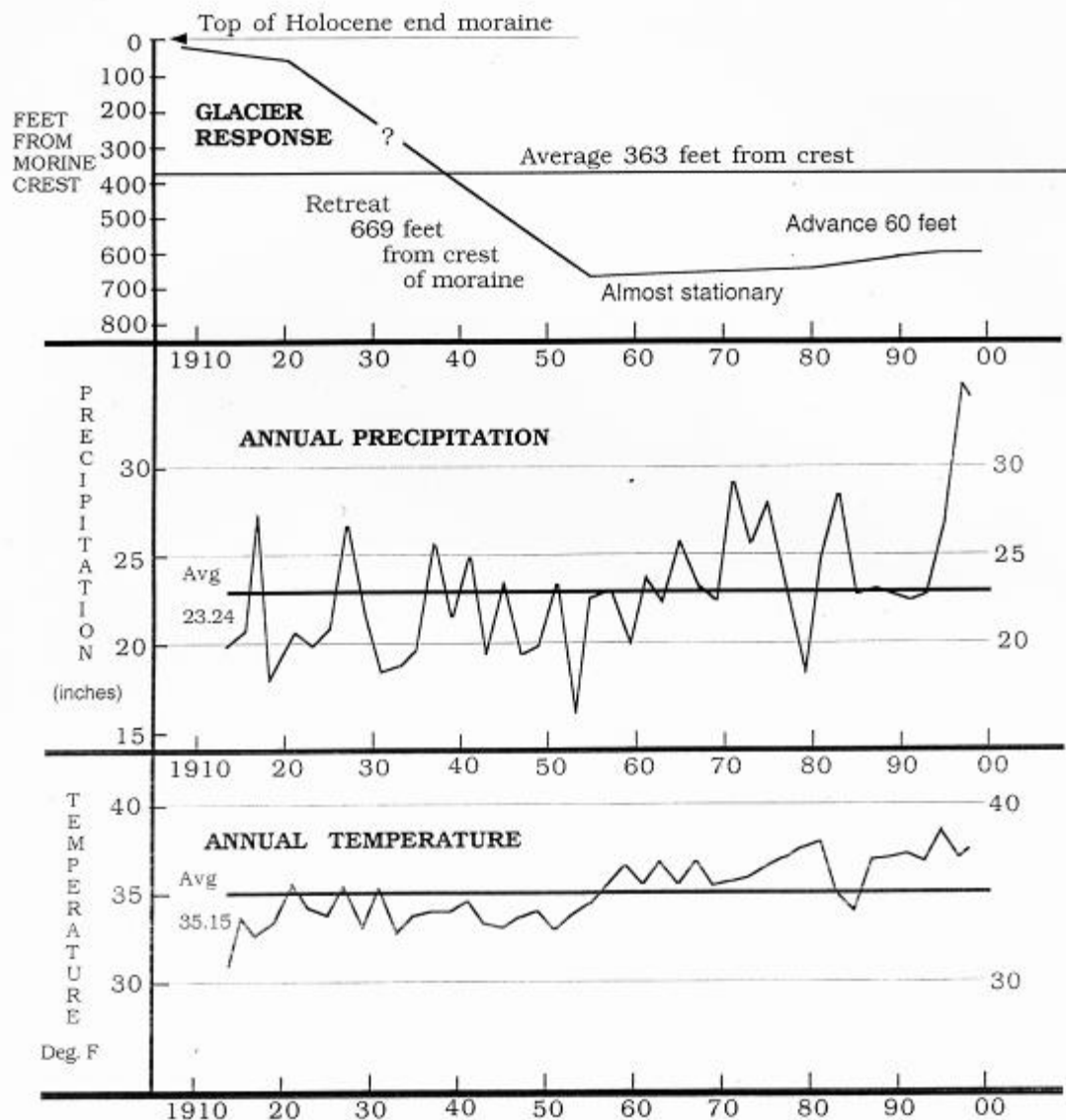
Cascade Canyon

Mount Saint John

Jenny Lake



TETON GLACIER RESPONSE TO CLIMATE CHANGES



RESPONSE OF TETON GLACIER TO CLIMATE CHANGES 1914 - 1998.
Weather records are from the Moran Weather Station in Teton National Park.

Average Precipitation 1914-1955 = 21.3 in. Average precipitation 1956 - 1998 = 24.5 in.
Average Temperature 1914-1955 = 34 deg F. Average Temperature 1956-1998 = 37.8 deg

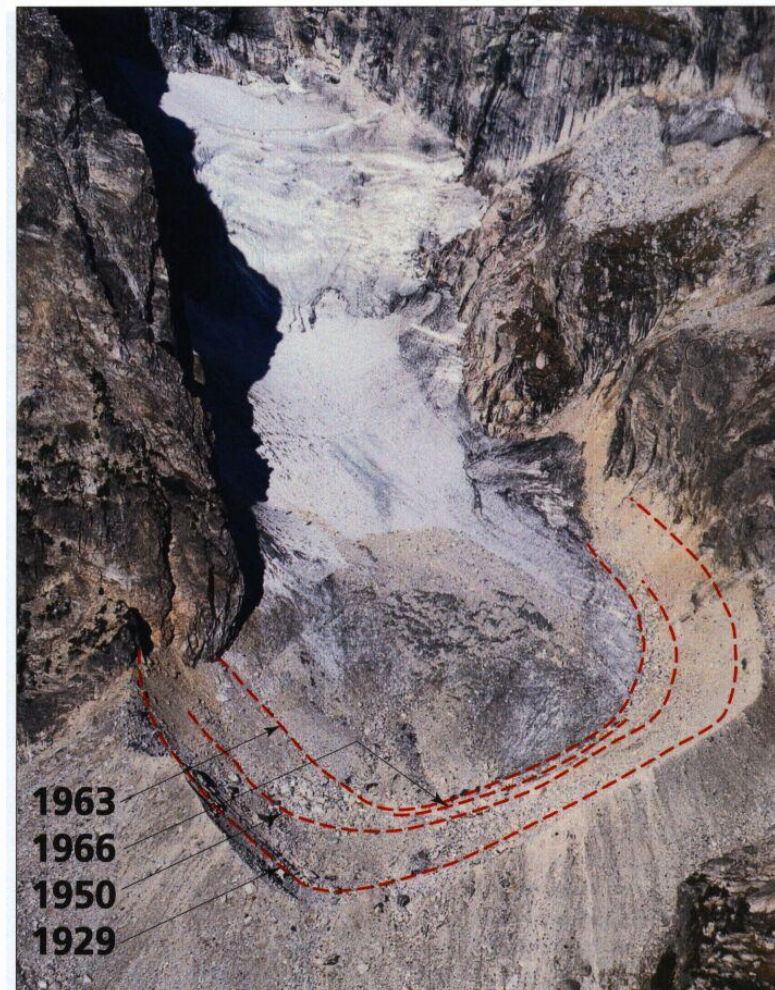
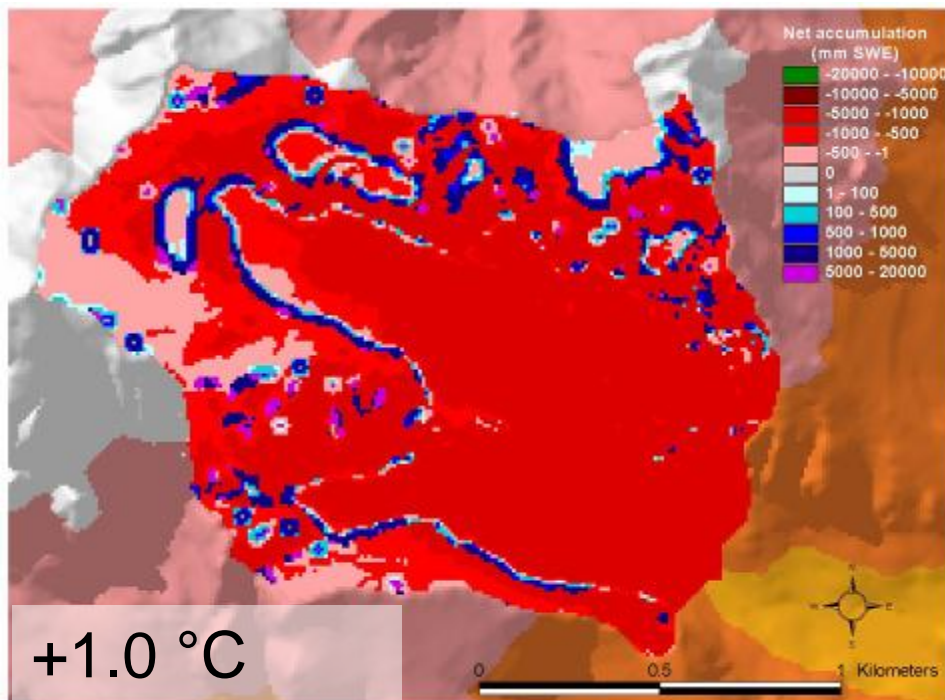
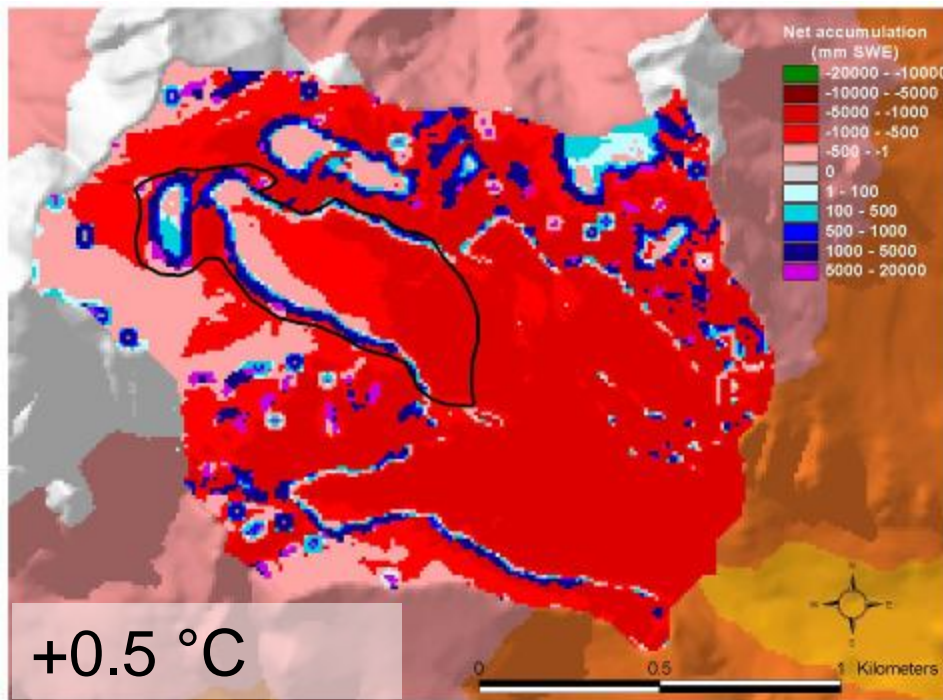
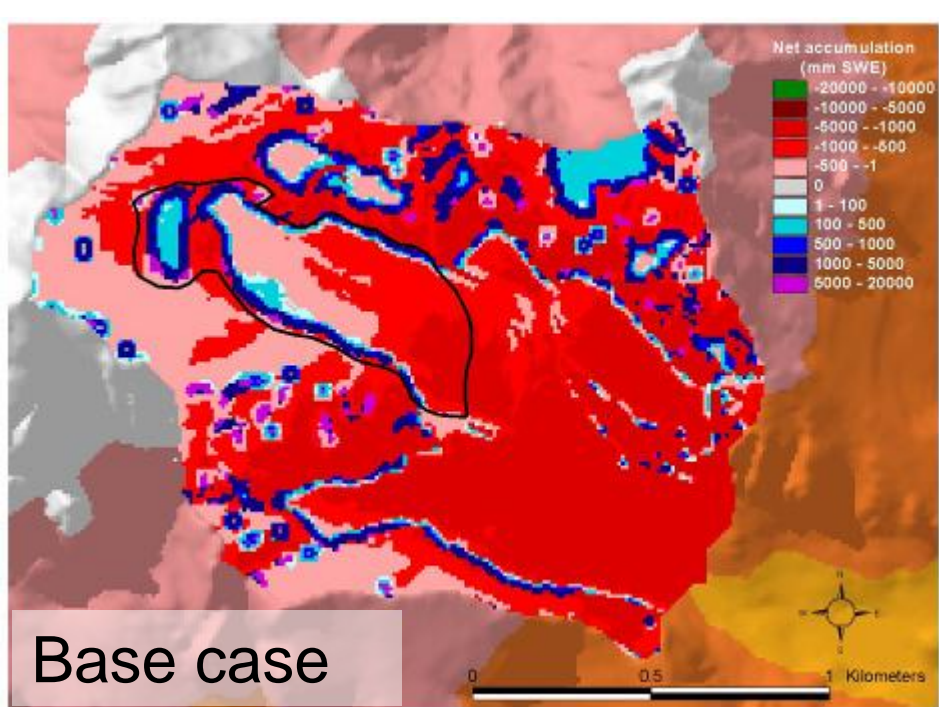
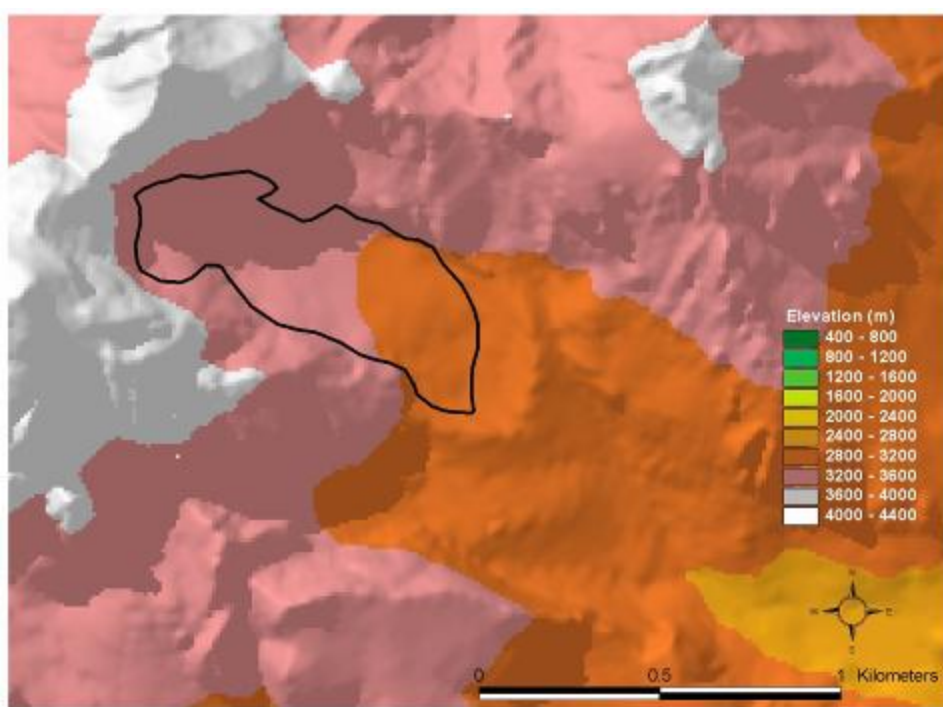
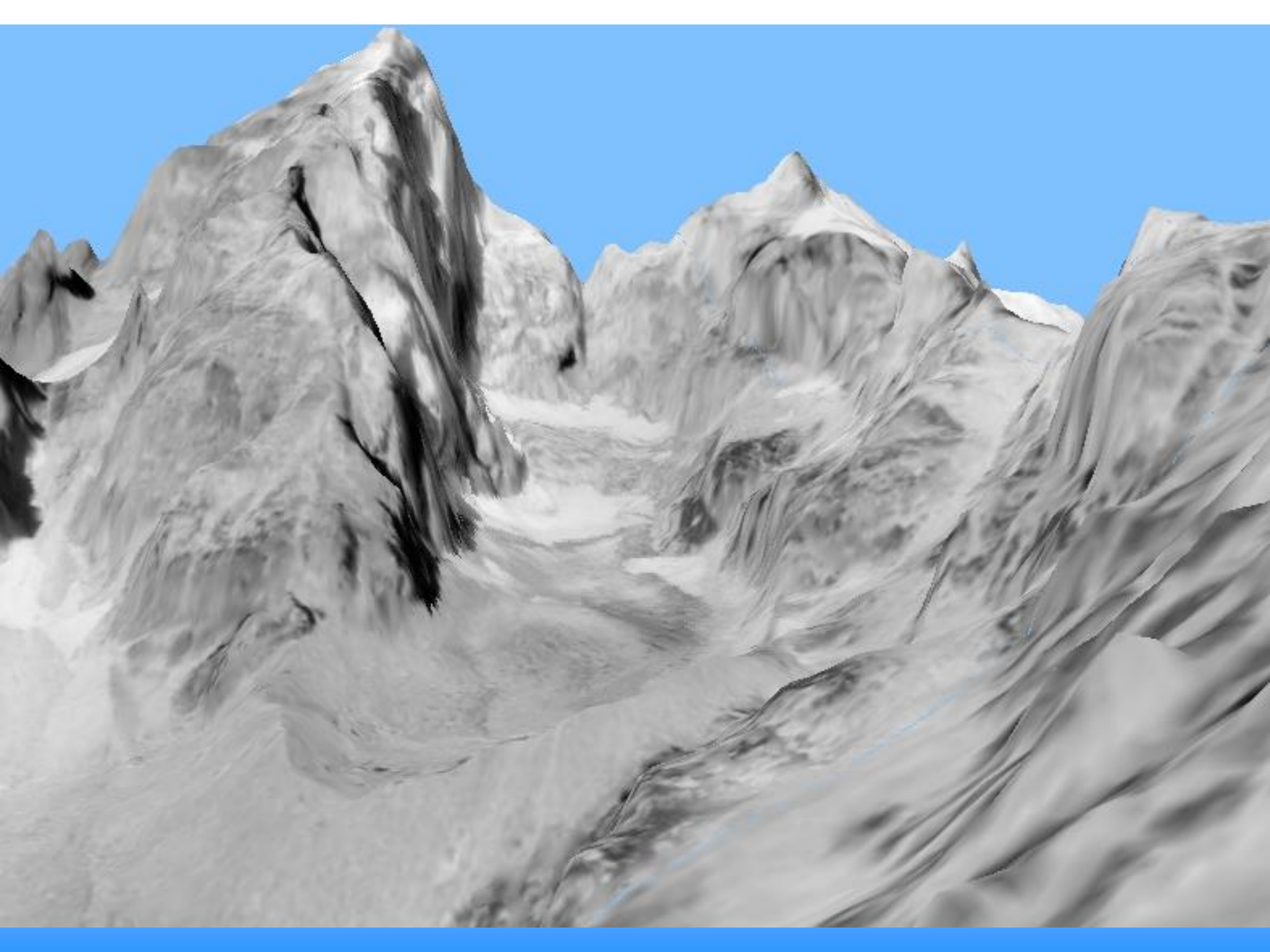
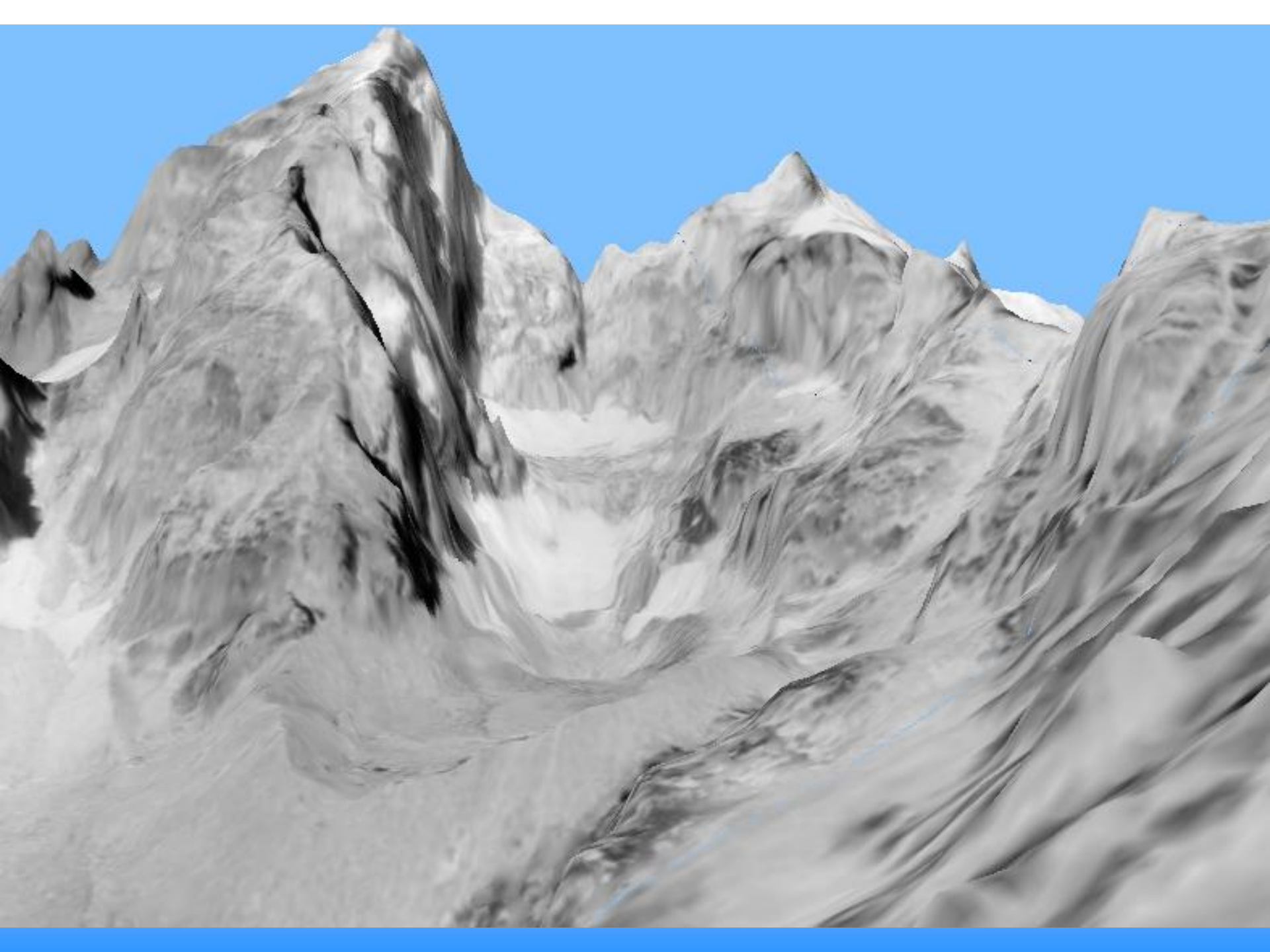


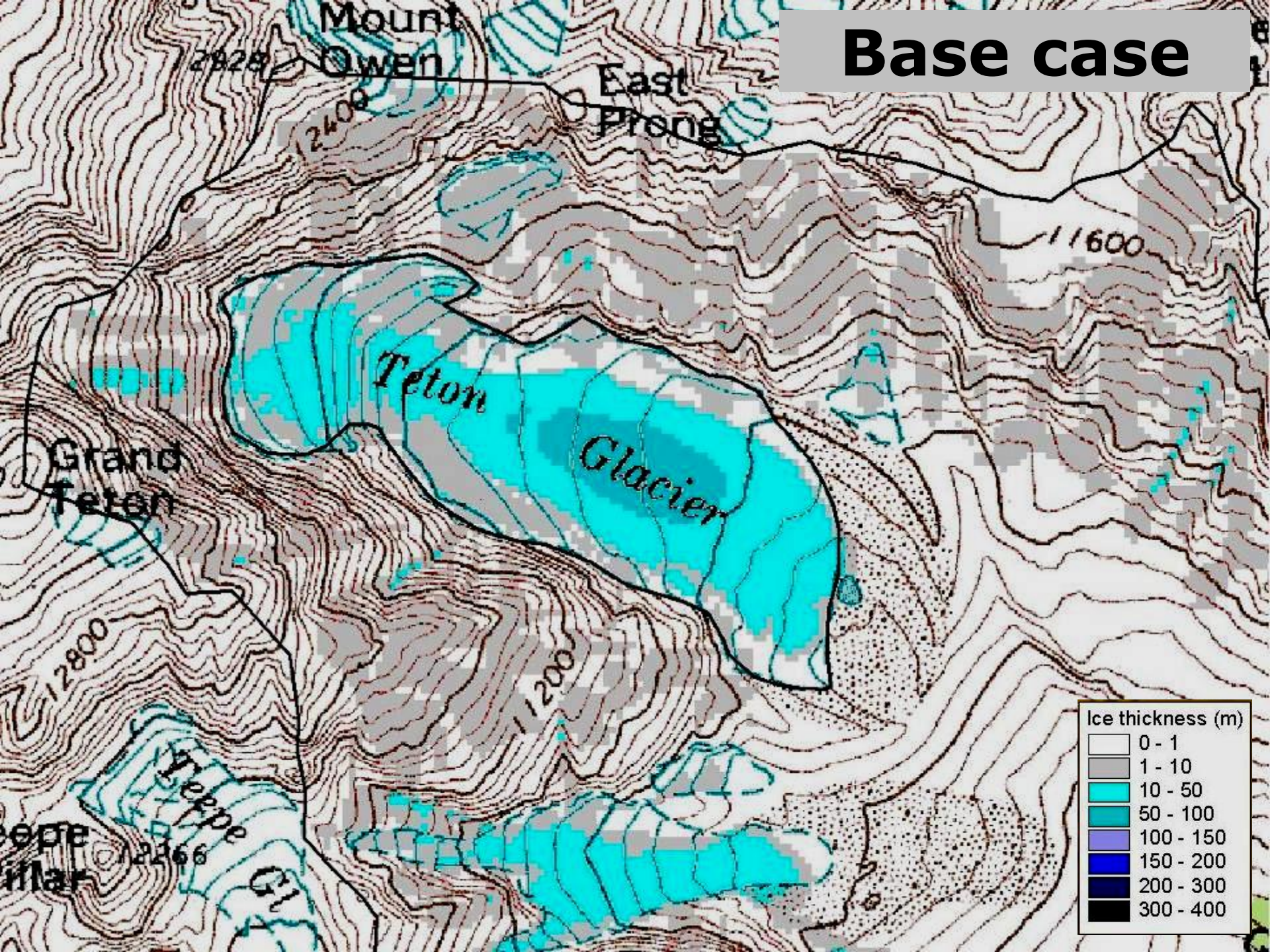
Figure 96: The Teton glacier, showing its retreat since its maximum advance during the Little Ice Age, represented by the sharp-crested fresh moraine. Approximate positions of the ice terminus at different times shown by red lines. Photo by E.J. Williams, 1998.



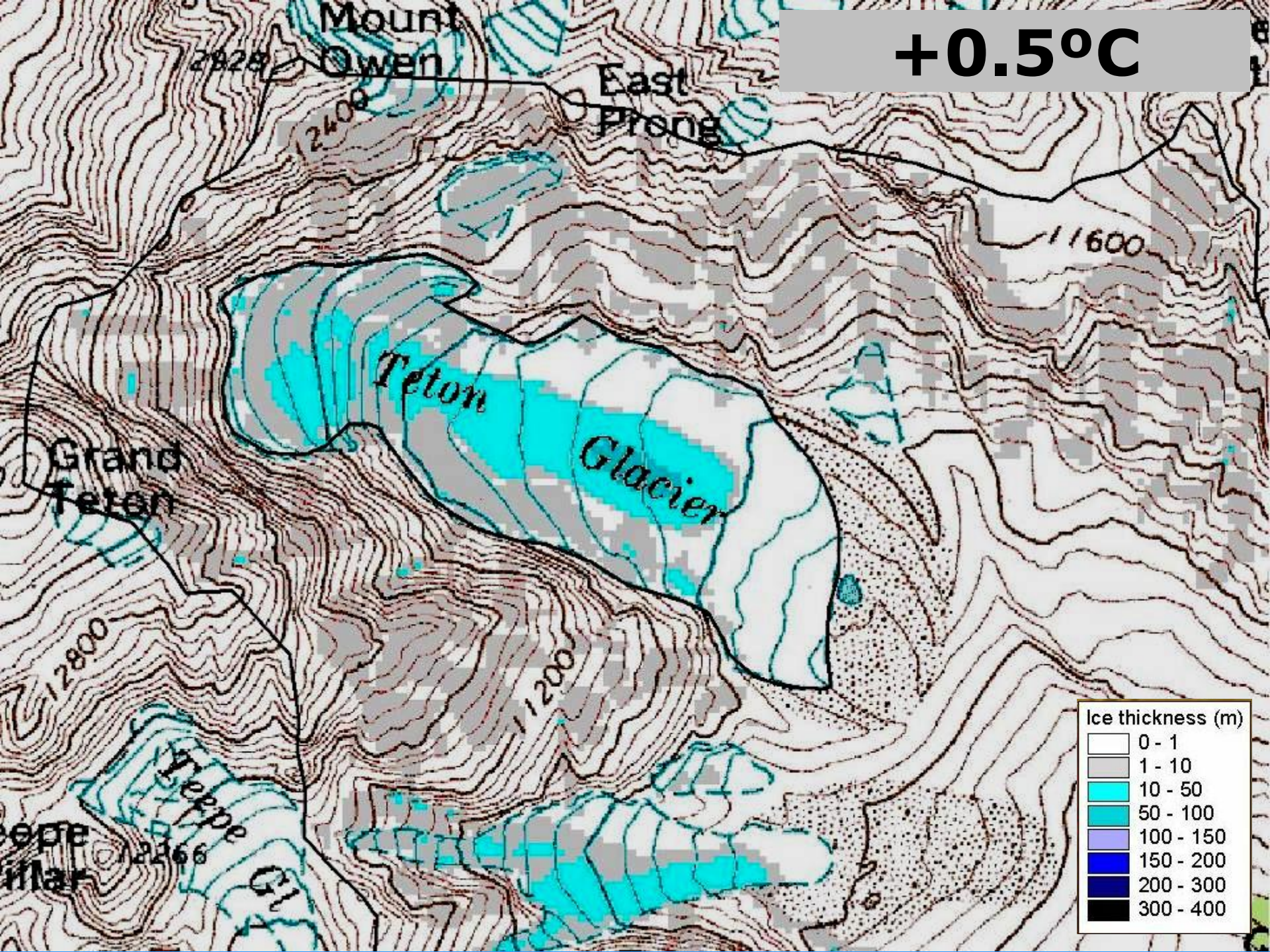




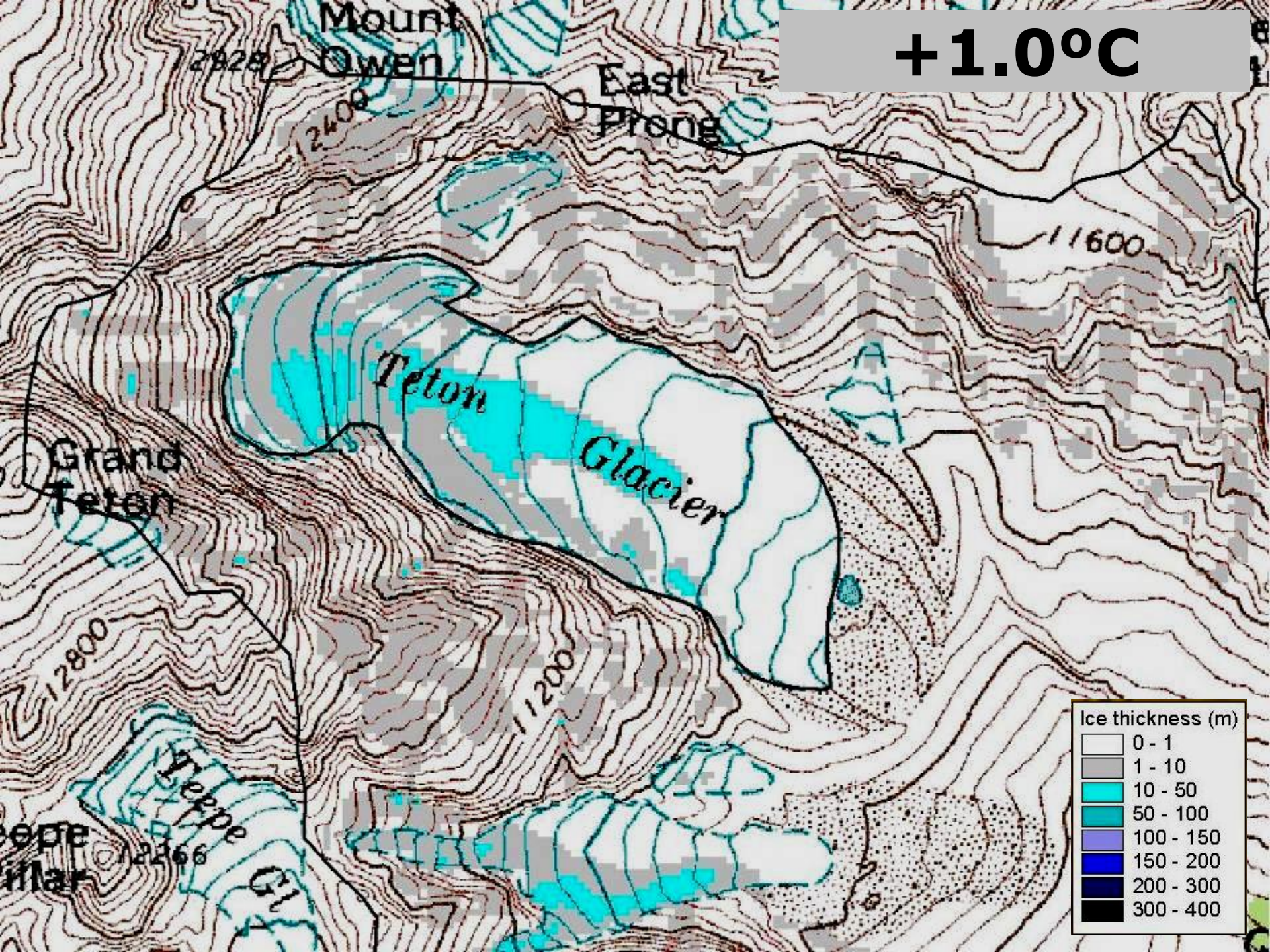
Base case



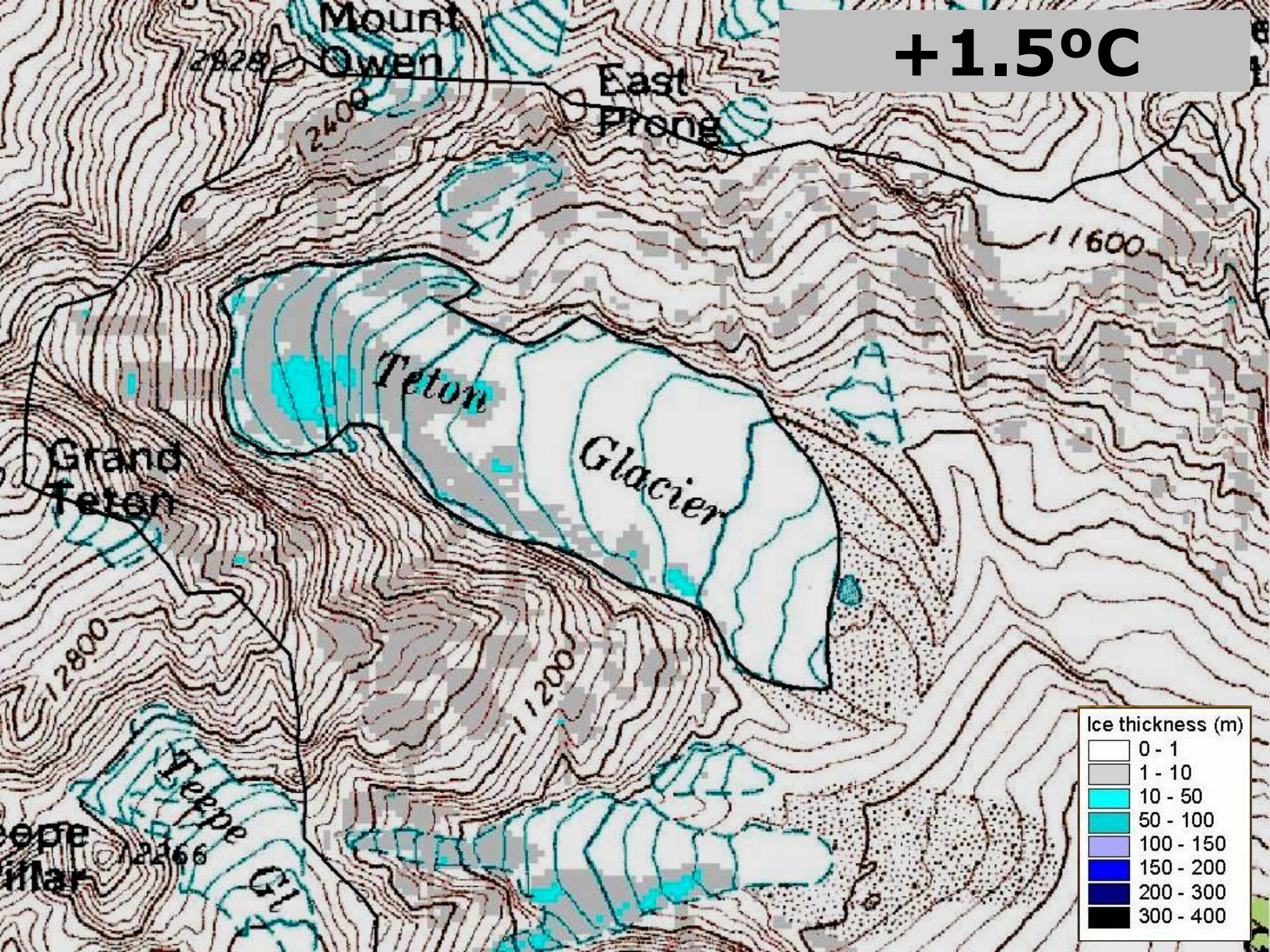
+0.5°C



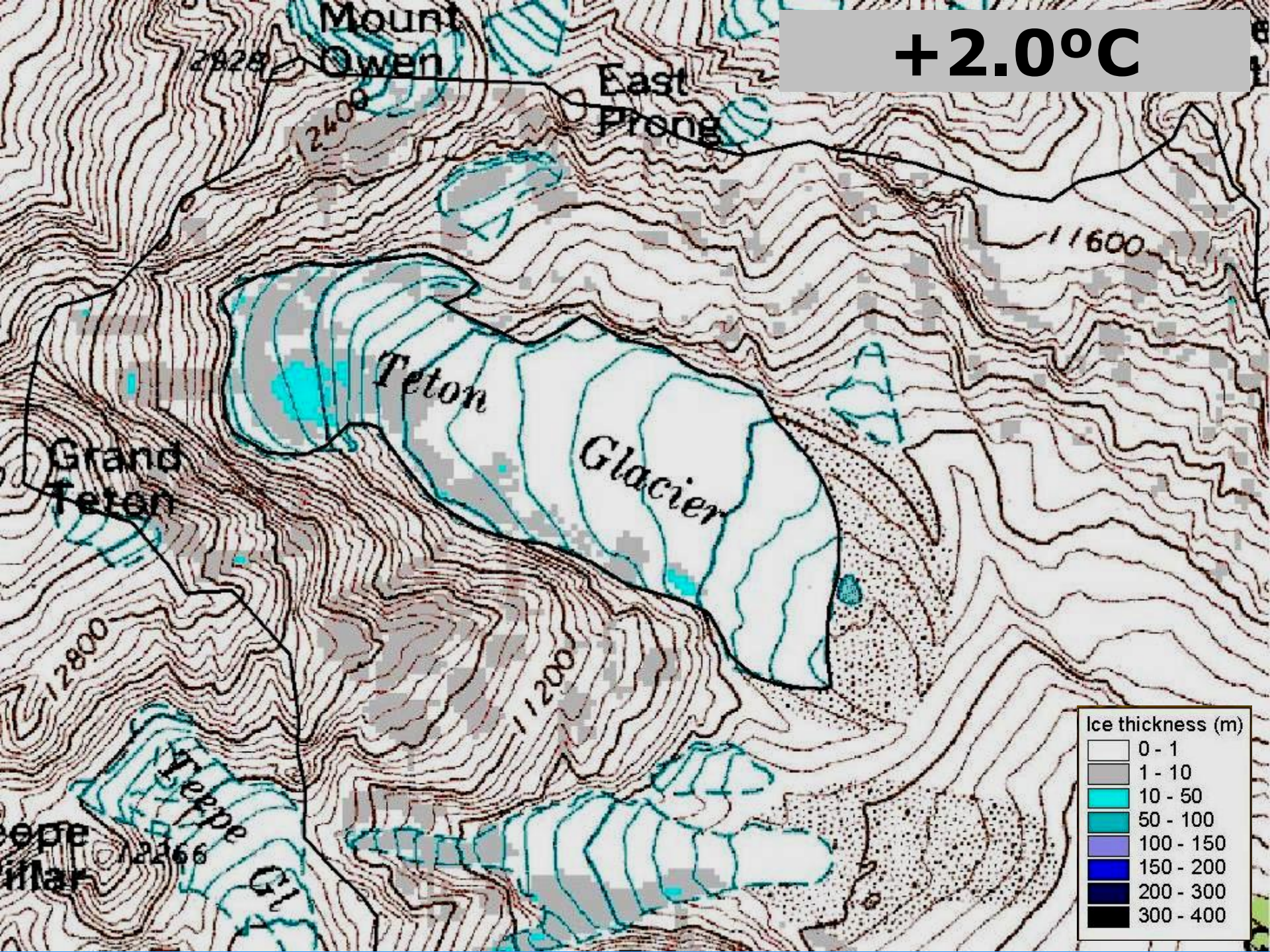
+1.0°C



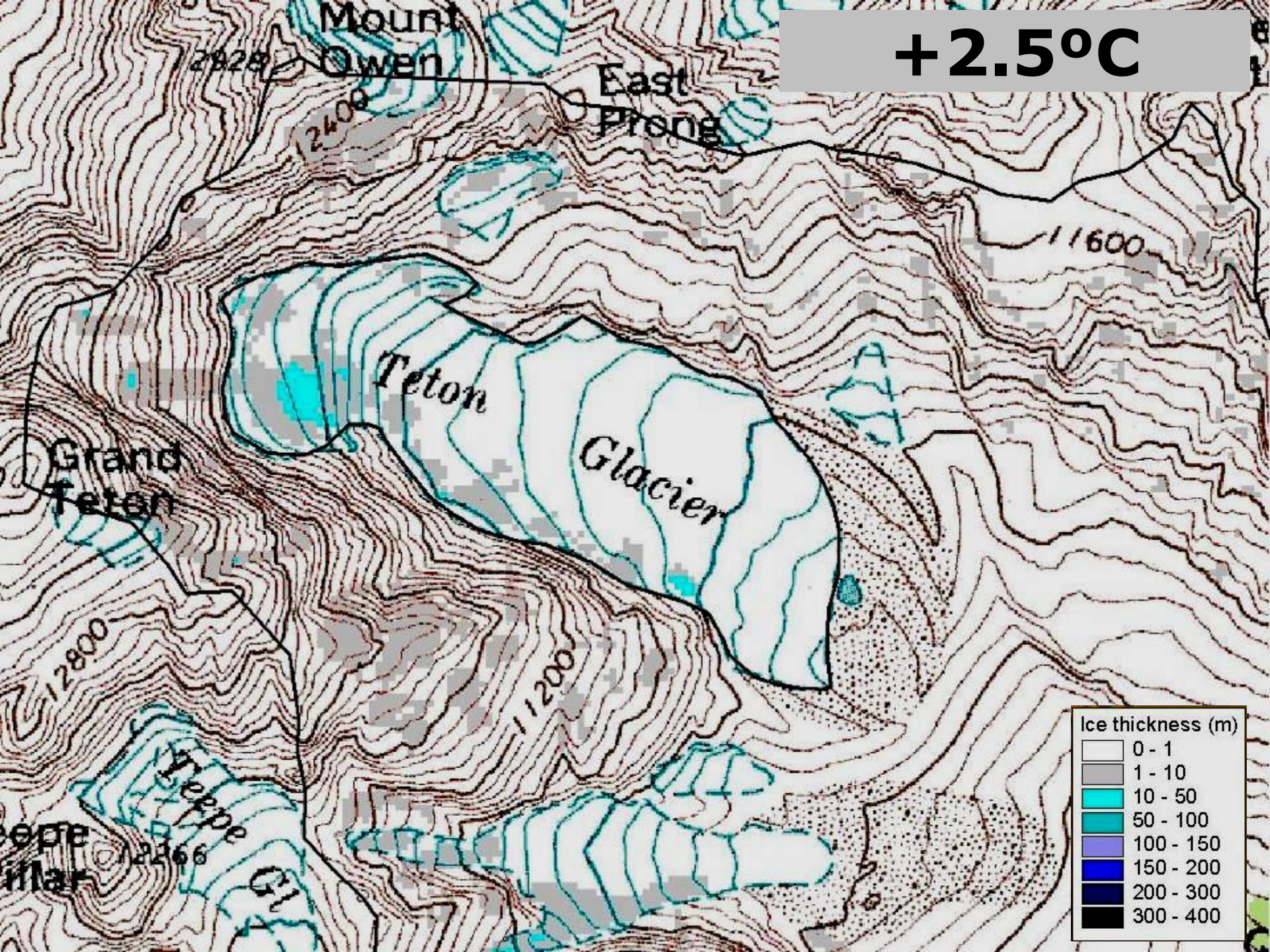
+1.5°C



+2.0°C



+2.5°C



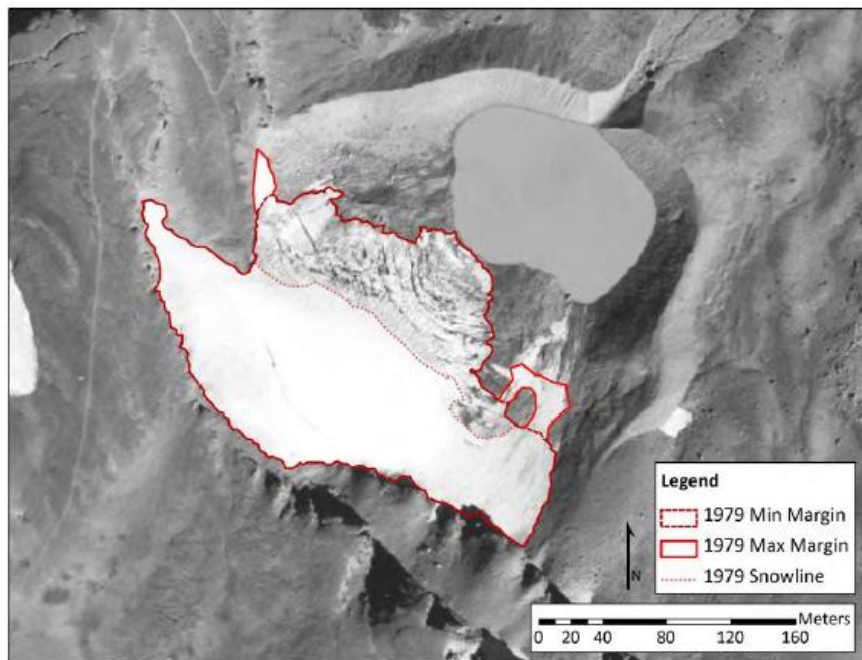


Figure 9. Schoolroom Glacier in 1979.

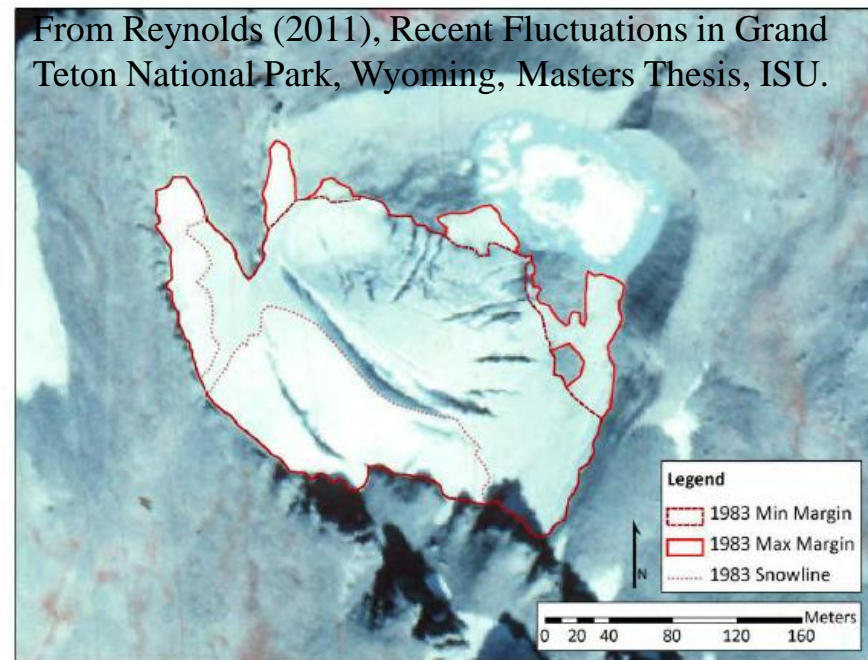


Figure 10. Schoolroom Glacier in 1983.

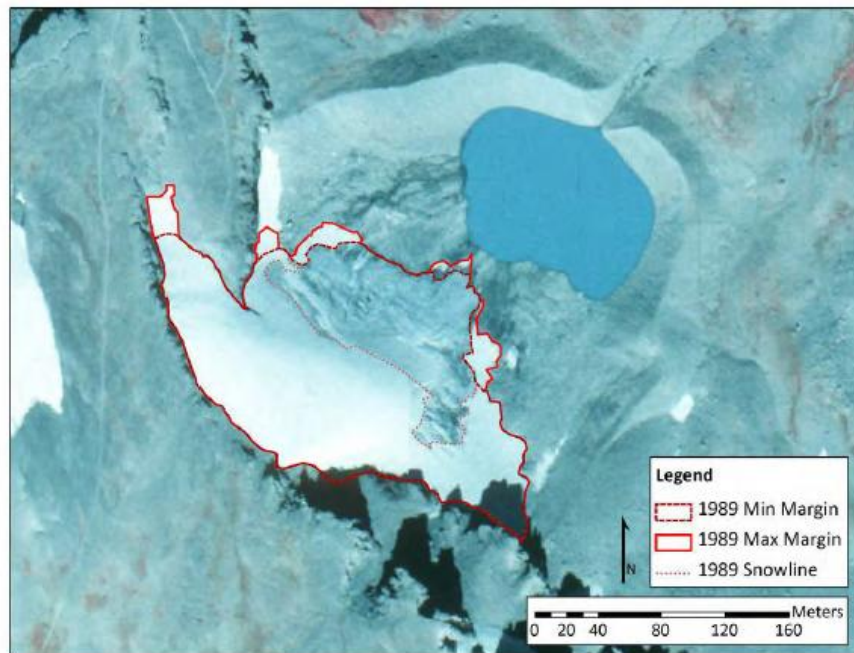


Figure 11. Schoolroom Glacier in 1989.

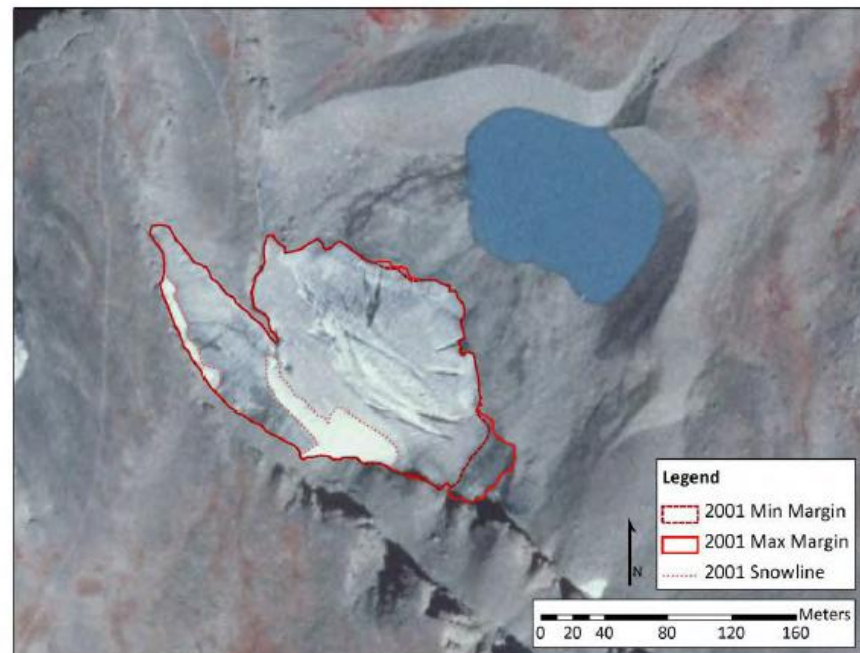
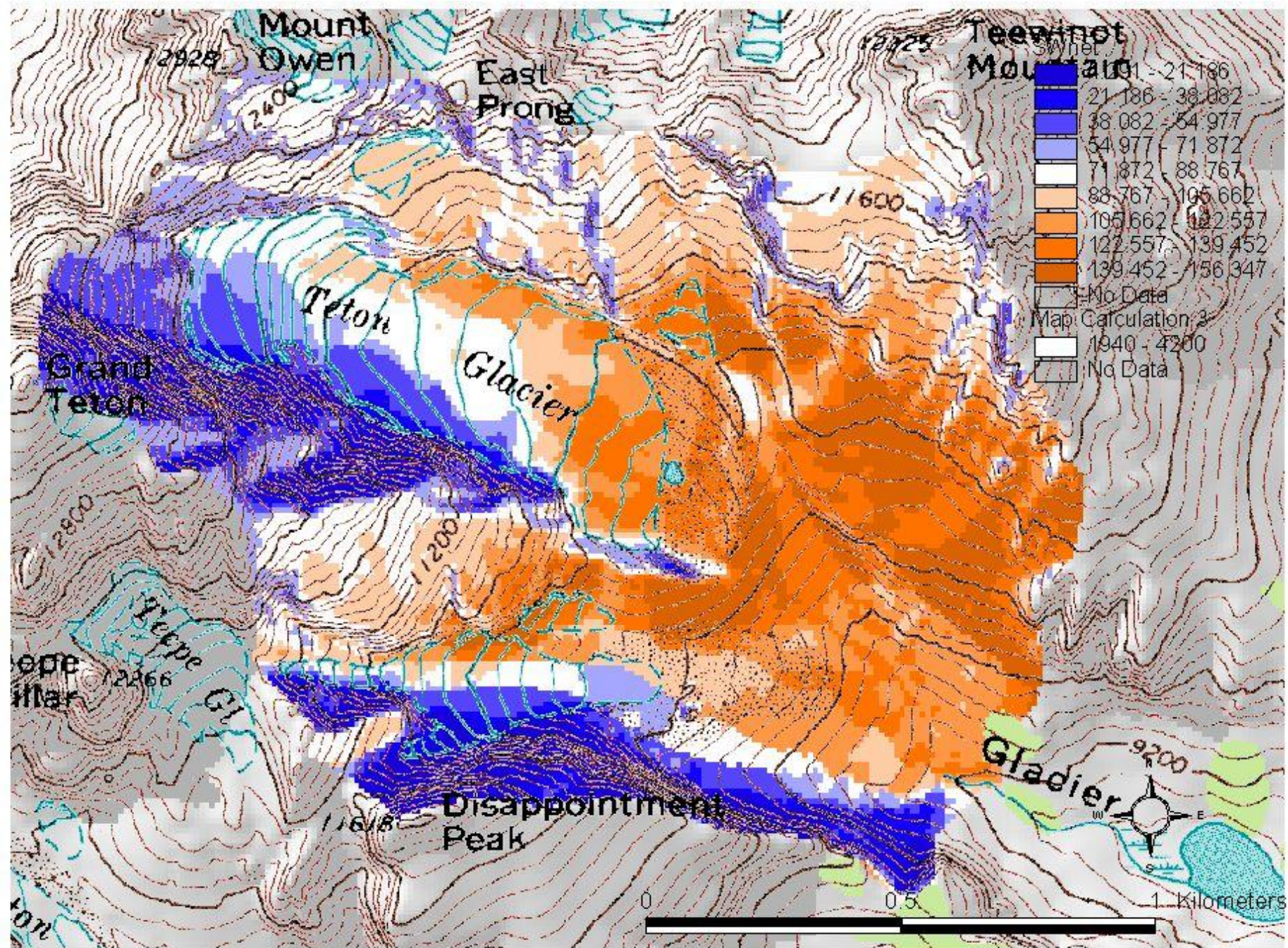


Figure 13. Schoolroom Glacier in 2001.

Remote glaciers tend to receive less attention than more easily accessible glaciers, even though they may be more strongly affected by climate warming (Arendt et al., 2002). Area fluctuation of glaciers in GTNP can be indicators not only of regional climate change, but also of the broader impacts of climate change on the region.



Skillet and Falling Ice Glaciers:

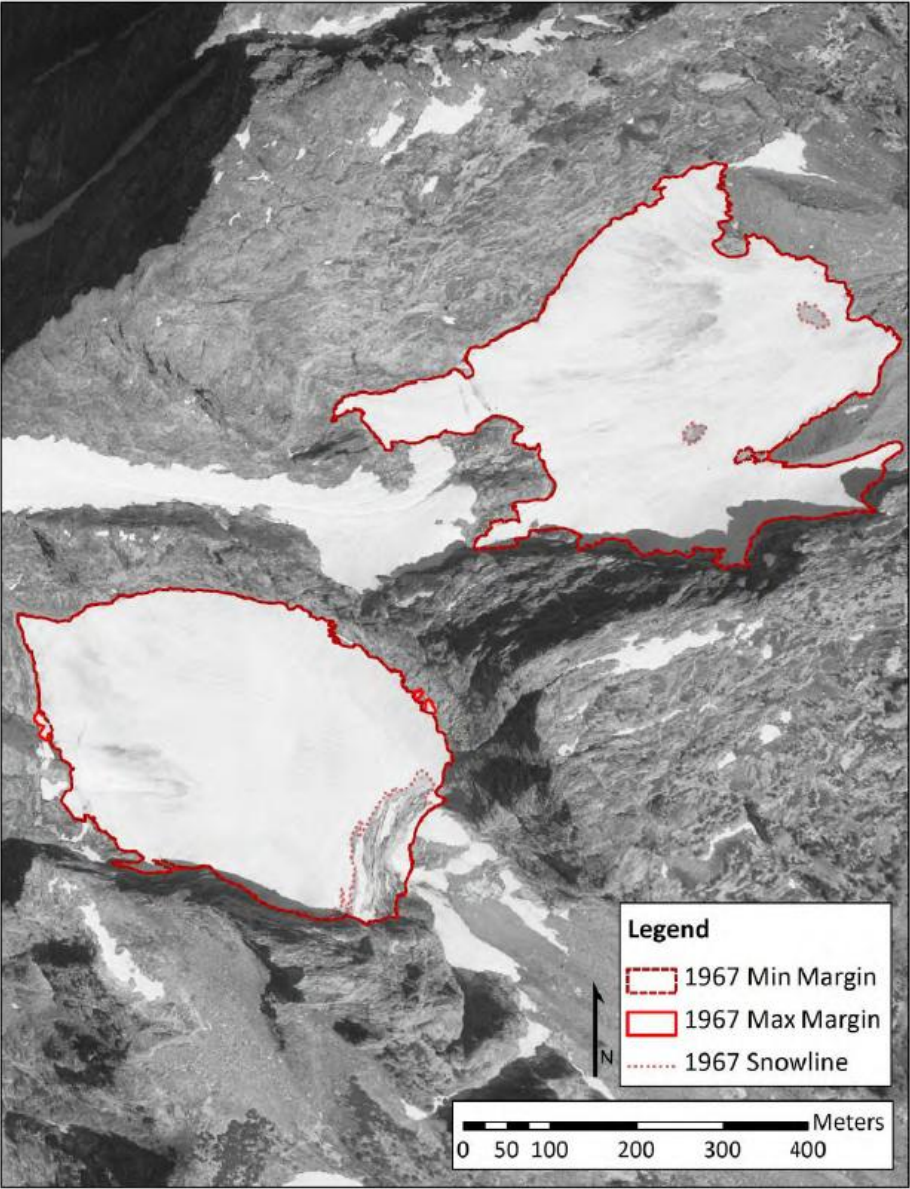


Figure 16. Skillet and Falling Ice Glaciers in 1967.

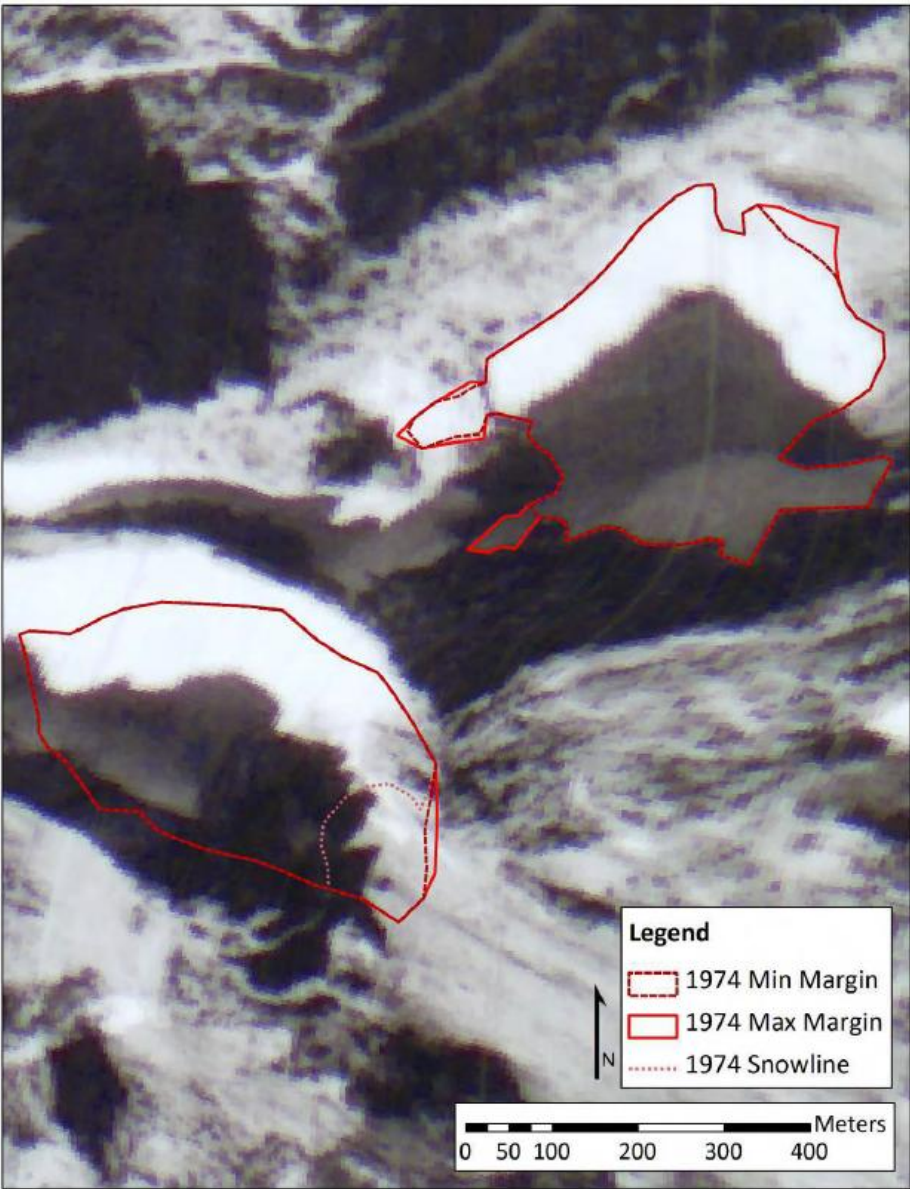


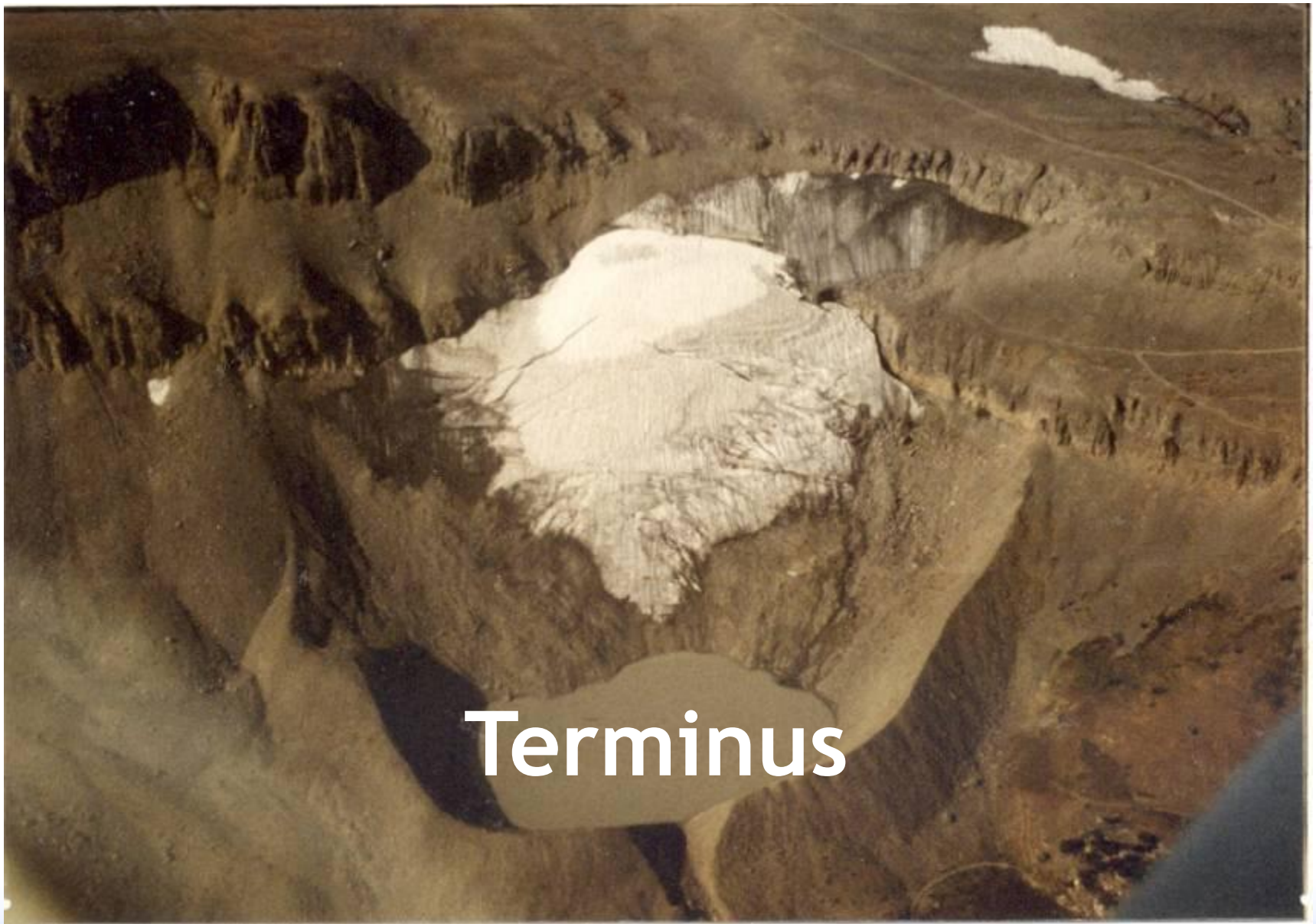
Figure 17. Skillet Glacier in 1974.

Falling Ice Glacier, Mt. Moran

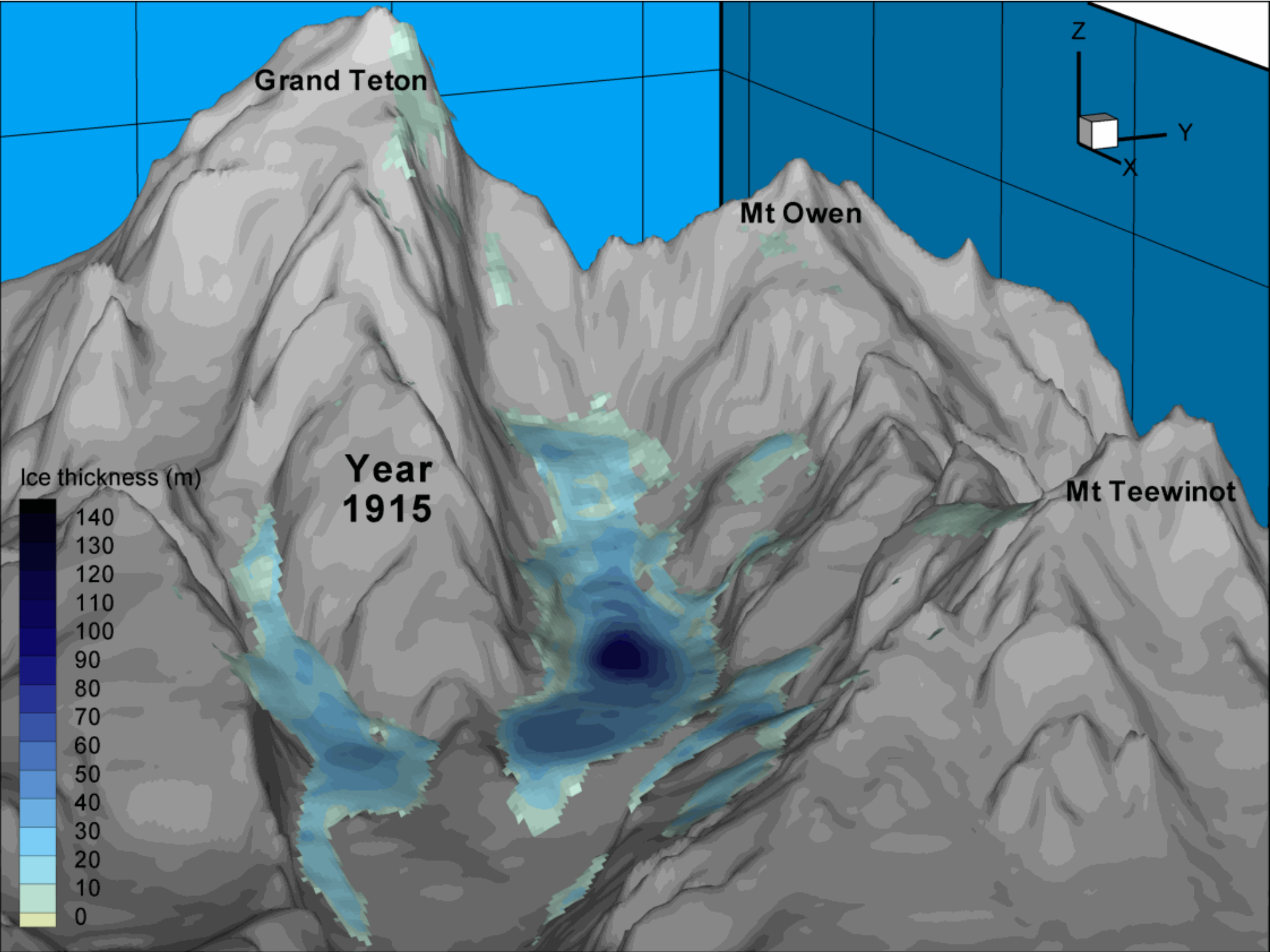


Conclusions

- Modeling experiments suggest that a temperature increase of $\sim 1.5^{\circ}\text{C}$ would dramatically reduce the volume of the Teton glacier
- At late 20th century global warming rates ($\sim 0.1^{\circ}\text{C}$ / decade), such retreat could take as long as 150 – 200 years
- Global climate models scaled to the Pacific NW project an increase in average temperature on the order of $0.2^{\circ}\text{-}0.6^{\circ}\text{C}$ ($0.3^{\circ}\text{-}1^{\circ}\text{F}$) per decade throughout the 21st century, implying a much faster recession.
- The mass balance profile of small alpine glaciers is often very steep near the equilibrium line altitude, suggesting decreased sensitivity relative to larger, more exposed, glaciers.
 - Future simulation efforts are aimed at improving resolution of model, in order to measure sensitivity of small alpine glaciers as they recede into increasingly shaded environments.



Terminus



Questions

- Barring 'local' redistribution of snow by wind, avalanching etc, how variable is precipitation across short distances in the mountains and how would one capture that in a simple model?
- In paleoglacier modeling,
 - How is a large change in mountain precipitation distributed vertically?
 - How reasonable is it to assume constant lapse rates?
 - Are pseudo-vertical lapse rates related to vertical precipitation gradients?
- Need more data to describe mountain climate and its variability